

# Inferring global-scale spatio-temporal $\delta^{18}O_{p}$ patterns from local datasets

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## **Background & Aims**

The abundance of publicly-available precipitation  $\delta^{18}O(\delta^{18}O_{p})$  data allows analysis of **global-scale**  $\delta^{18}$ **O**, **patterns**. Most such meta-analyses focus on static **spatial**  $\delta^{18}O_{p}$  patterns.

Characterising global-scale  $\delta^{18}O_{p}$  variability **through time** is more difficult.  $\delta^{18}O_{p}$ data availability is spatially & temporally irregular, as shown on the plot below which summarises the temporal coverage of publicly-available  $\delta^{18}O_{p}$  station data.



Here we aim to characterise both the spatial and temporal variability in **global**  $\delta^{18}O_{P}$ , using data from GNIP and the Water Isotopes Database.

### **Methods**

We grouped  $\delta^{18}O_{p}$  stations into clusters based on geographic & climatic parameters. The map to the right shows  $\delta^{18}O_{\rm p}$  stations coloured according to their assigned cluster (52 clusters total).



We used a novel 'dynamic compositing' method to combine all records in each cluster into a single timeseries, without spurious jumps in mean or variance.



From those 52 clusters we kept only the **16 regional**  $\delta^{18}O_{\rm p}$  composites with >80 % temporal coverage from 1982 to 2015.

We found the common gradient underlying these regional  $\delta^{18}O_{p}$  composites (global  $\delta^{18}O_{p}$  PC1 & EOF1, at annual resolution), and did the same on simulations from the isotope-enabled CESM for comparison.

We then compared the global  $\delta^{18}O_{\rm p}$  PC1 with globally-relevant climate indices.

## **Results**



Orange: positive loading on PC1. Purple: negative loading on PC1

The three maps below show the correlation of global  $\delta^{18}O_{p}$  PC1 with climatic variables (sea level pressure, precipitation, and near-surface winds).







Purple & orange **points** on the map to the left show the loading of regional  $\delta^{18}O_{p}$  composites on the global  $\delta^{18}O_{p}$  PC1 i.e. as PC1 increases,  $\delta^{18}O_{\rm p}$  at purple-coloured sites decreases &  $\delta^{18}O_{\rm p}$  at orange-coloured sites increases.

Small grey circles show the individual sites comprising the regional  $\delta^{18}O_{P}$  composites.

The **background grid** shows  $\delta^{18}O_{p}$ EOF1 from the isotope-enabled CESM.

Precipitation amount

### Discussion

The table below shows the **correlation** of global  $\delta^{18}O_{P}$  PC1 with globally-relevant climatic indices

Tropical Pacific atmospheric variability	PWC (ΔSLP)	SOI	
	0.74*	0.70*	
Tropical Pacific oceanic variability	Niño 3.4	EMI	
	-0.58*	-0.44*	
Variability outside the tropical Pacific	PMM	SAM	DMI
	-0.33	0.24	-0.03

PWC = Pacific Walker Circulation (as defined by trans-Pacific SLP gradient), SOI = Southern Oscillation Index, EMI = ENSO-Modoki Index, PMM = Pacific Meridional Mode, SAM = Southern Annular Mode, DMI = Dipole Mode Index. Asterisk denotes significant correlation

Global  $\delta^{18}O_{P}$  PC1 is most strongly correlated with the strength of the **Pacific Walker** Circulation.

### We see similar isotope-climate relationships in observations & the iCESM, suggesting both that iCESM accurately models $\delta^{18}O_{\rm p}$ patterns, & that our 16 regional $\delta^{18}O_{\rm p}$ composites capture a **realistic approximation** of spatio-temporal changes in global $\delta^{18}O_{p}$ .

Our methodology could also be used to synthesise regional-scale  $\delta^{18}O_{P}$  variability from local  $\delta^{18}O_{p}$  data, or to analyse  $\delta^{18}O_{p}$  variability in different time periods.

For all the details, see our recent paper (open access): https://tinyurl.com/falster2021

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