

Kinetic fractionation factors for ocean evaporation: limited impact of wind speed observed

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Isotope non-equilibrium fractionation effects are important during ocean evaporation and sensitive to environmental conditions such as temperature and relative humidity in the lower troposphere.

The Craig-Gordon model can be used to describe the fractionation effects involved during the evaporation process. The model can be represented as a **multi-layer model** where different isotopic effects are involved.

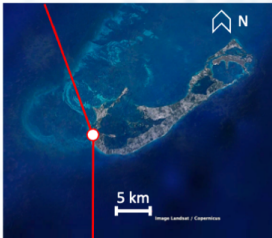
$$\delta_E = (1 - k) \frac{\alpha_V(1 + \delta_L) - h(1 + \delta_A)}{(1 - h)} - 1$$

The kinetic effects are summarized by the *k* value, hereafter:

$$k = \frac{\left(\frac{D}{D_i}\right)^n - 1}{\left(\frac{D}{D_i}\right)^n + \frac{p_T}{\rho_M}}$$

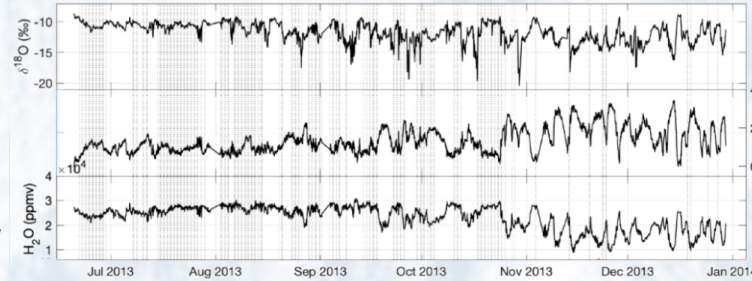
ρ/ρ_M can be expressed as a function of the **wind velocity vertical distribution** and ultimately to wind speed at a reference height (e.g. 10 m).

The objectives of this study are to estimate the **best *k* values for ocean evaporation** and to quantify the **wind speed effect** on kinetic fractionation in oceanic conditions



We used six months of water vapor isotopic observations collected from a meteorological tower located in the northwest Atlantic (Tudor Hill Marine Atmospheric Observatory - THMAO, Bermuda).

At THMAO, water vapor isotopic composition was measured at two heights between June to December 2013.



The **isotopic composition of evaporation flux** was calculated with the **Craig Gordon model** and estimated from 2-inlets observations at THMAO using **Flux Gradient and Keeling Plot** method.

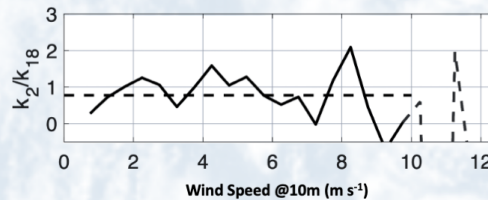
The **observed flux composition** was compared to **calculated flux composition**. This way, best kinetic fractionation values for D/H and ¹⁸O/¹⁶O were estimated by **minimization of the differences between model and observations**.

The mean *k*-values estimated by error minimization are: $k_{18} = 5.2 \pm 0.6 \%$ and $k_2 = 4.3 \pm 3.4 \%$

The best-fit line between k_{18} and 10-m wind speed ($m s^{-1}$) is:

$$k_{18} = (-0.19 \pm 0.04) * WS + (6.7 \pm 0.3) \%, R^2=0.52$$

The **sensitivity of k_{18} to WS is only $-0.19 \pm 0.06 \%$ $m^{-1}s$** . k_2 observations are more affected by noise ($R^2=0.11$). The linear relationship between *k* and WS is more robust for k_{18} .

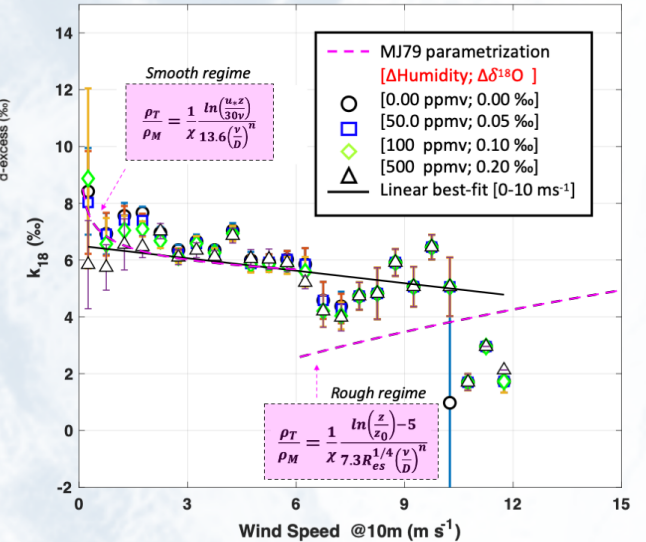


The k_2/k_{18} ratio shows **no significant correlation with WS**. Between 0.5 and 10 $m s^{-1}$, the k_2/k_{18} ratio largely varies around the mean value of 0.8 with a standard deviation of 0.6.

$$k_2/k_{18} = (0.8 \pm 0.6) * k_{18}$$

Modulating the threshold of differences between water vapor observations at the two inlets affects low-wind speed observations

Wind speed sensitivity is higher at low wind speed, consistent with MJ79 smooth parametrization



This study provided **evidence of wind speed effect in modulation of kinetic fractionation factor** during ocean evaporation. This effect is very small but statistically significant.

This study **do not provide enough evidence of different wind speed regimes** (e.g. smooth and rough) modulating the kinetic fractionation effect

